

Summarised description

Unit name

Jupille Member

Code

JUP

Status

Formal Member

Parent unit

La Roche Formation

Child units

None

Characteristic description

Slaty siltstone and shale with sandstone intercalations.

Age

Pragian

Thickness

0-600 metres

Area of occurrence

Southern and south-eastern limbs of the Dinant Synclinorium east of the Lomme River valley and in the Neufchâteau-Eifel Synclinorium east of La Roche-en-Ardenne.

Type locality

Sections along the road N833 from La Roche-en-Ardenne to Hotton, c. 600 m south of Jupille (Dejonghe et al., 2008).

Alternative names

None

Authors

Dejonghe et al. (2008): *Formation de Jupille*

Modified after

Denayer, J. & Mottequin, B., 2024. Lower Devonian lithostratigraphy of Belgium. *Geologica Belgica*, 27/3-4, 115–154.

Date

23/09/2025

Cite as

Denayer, J. & Mottequin, B., 2025. The Jupille Member, 23/09/2025. National Commission for Stratigraphy Belgium. <https://ncs.naturalsciences.be/lithostratigraphy/Jupille-Member>

Full description

Unit name

Jupille Member

Code

JUP

Status

Formal Member

Parent unit

La Roche Formation

Child units

None

Origin of the name

After the town of Jupille in the Ourthe valley.

Alternative names

None

Authors

Dejonghe et al. (2008): *Formation de Jupille*

Modified after

Denayer, J. & Mottequin, B., 2024. Lower Devonian lithostratigraphy of Belgium. *Geologica Belgica*, 27/3-4, 115–154.

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Characteristic description

In the Ourthe River valley, this unit is better individualised and was described as the Jupille Formation by Dejonghe et al. (2008). However, the impossibility to pinpoint a boundary between the La Roche and the Jupille formations, as well as the very variable development of sandstone make it complicated to separate both lithostratigraphic units. Therefore, we here chose to retrograde the sandy unit to the member level, which is also more concordant with the previous studies (Asselberghs, 1946; Stainier, 1994). In its stratotype, the Jupille Member is composed of decimetric beds of dark grey to brownish sandstone and siltstone alternating with dark micaceous shale. The sandstone and siltstone beds, which are locally micaceous and bioturbated, display a wide range of sedimentary structures (oblique stratifications, toolmarks, ripples, lenticular bedding, oblique and cross stratifications, load casts, etc.). They typically have a brownish-orange weathering colour due to their limonitic content. In the upper part of the Member, the sandstone is slightly arkosic and displays ball-and-pillow structures. Eastwards (Nisramont, Lengeler), the proportion of sandstone decreases, and the dark shale (and slate locally) becomes dominant.

Area of occurrence

Southern and south-eastern limbs of the Dinant Synclinorium east of the Lomme River valley and in the Neufchâteau-Eifel Synclinorium east of La Roche-en-Ardenne.

Type locality

Discontinuous section along the road N833 from La Roche-en-Ardenne to Hotton, c. 600 m south of Jupille (Dejonghe et al., 2008).

Age

Mostly late Pragian, after the age of the La Roche Formation (Godefroid & Stainier, 1982).

Thickness

Highly variable: c. 600 m in Jupille and Nisramont (Dejonghe et al., 2008) but only 250 m in Lengeler and decreases to 0 m in the Grand Duchy of Luxembourg (Dejonghe, 2019).

Lower boundary

The base of the member is not easy to define, as the appearance of sandstone beds is transitional.

Upper boundary

The increasing proportion of the sandstone bed in the Jupille Member marks the transition to the overlying Ruisseau de la Forge Formation, though the contact has never been observed.

Regional correlations

In the north-eastern part of the Neufchâteau–Eifel Synclinorium, the La Roche Formation passes to the monotonous slaty facies of the Kautenbach–Troisvierges Formation—the *facies de Saint-Vith* sensu Asselberghs (1927, p. 209; 1946, p. 179)—named *Wüstebach-Formation* on the German territory (Ribbert et al., 1992). In the Neufchâteau–Eifel Synclinorium, the upper part of the La Roche Formation tends to be richer in sandstone and quartzitic sandstone beds, which probably corresponds to the transition to the overlying Our Formation (Brichant, 1928; Brül, 1966).

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